



1 **Ocean acidification indirectly alters trophic interaction of**
2 **heterotrophic bacteria at low nutrient conditions**

3

4 **Thomas Hornick¹, Lennart T. Bach², Katharine J. Crawford³, Kristian Spilling^{4,5},**
5 **Eric P. Achterberg^{2,6}, Corina P. D. Brussaard^{3,7}, Ulf Riebesell², Hans-Peter**
6 **Grossart^{1,8*}**

7

8 [1]{Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB), Experimental
9 Limnology, 16775 Stechlin, Germany}

10 [2]{GEOMAR Helmholtz Centre for Ocean Research Kiel, Düsternbrooker Weg 20, 24105
11 Kiel, Germany}

12 [3]{Department of Biological Oceanography, NIOZ – Royal Netherlands Institute for Sea
13 Research, P. O. Box 59, 1790 AB Den Burg, Texel, The Netherlands}

14 [4]{Marine Research Centre, Finnish Environment Institute, P.O. Box 140, 00251 Helsinki,
15 Finland}

16 [5]{Tvärminne Zoological Station, University of Helsinki, J. A. Palménin tie 260, 10900
17 Hanko, Finland}

18 [6]{National Oceanography Centre Southampton, European Way, University of Southampton,
19 Southampton, SO14 3ZH, UK}

20 [7]{Aquatic Microbiology, Institute for Biodiversity and Ecosystem Dynamics, University of
21 Amsterdam, P.O. Box 94248, 1090 GE Amsterdam, The Netherlands}

22 [8]{Potsdam University, Institute for Biochemistry and Biology, 14469 Potsdam,
23 Maulbeerallee 2, Germany}

24 Correspondence to: T. Hornick (hornick@igb-berlin.de)

25

26 **Abstract**

27 Annually, the oceans absorb about one fourth of the anthropogenically produced atmospheric
28 carbon dioxide (CO₂) resulting in a drop in surface water pH, a process termed ocean
29 acidification (OA). Surprisingly little is known about how OA affects physiology as well as
30 food web interactions of heterotrophic bacteria when essential nutrients are limited, since
31 most previous experiments were carried out during productive phases or even after nutrient
32 additions to stimulate algal blooms. Therefore, we conducted an *in situ* large-volume
33 mesocosm (~55 m³) experiment in the Baltic Sea by simulating different fugacities of CO₂
34 (*f*CO₂) extending from present to future conditions. The study was carried out after the spring-
35 bloom in July-August to maintain low-nutrient conditions throughout the experiment, which
36 resulted in a small-sized phytoplankton community dominated by picophytoplankton. Several
37 positive as well as negative effects on free-living (FL) and particle-associated (PA) bacterial
38 protein production (BPP) and biovolume (BV) could be related to *f*CO₂-induced differences in
39 phytoplankton composition and subsequent the availability of phytoplankton-derived organic
40 matter. However, dynamics of BV and cell-specific BPP (csBPP) of FL heterotrophic bacteria
41 could not be explained exclusively by the availability of phytoplankton-derived organic
42 carbon. The dynamics were also related to enhanced grazing on DNA rich (HDNA) bacterial
43 cells at higher *f*CO₂ as revealed by flow cytometry. Additionally, a decoupling of autotrophic
44 production and heterotrophic consumption during the last third of the experiment resulted in
45 low, but significantly higher accumulation of DOC at enhanced *f*CO₂. Interestingly we could
46 not detect any consistent and direct *f*CO₂-induced effect on BPP, csBPP nor BV of either FL
47 or PA heterotrophic bacteria. In contrast, our results reveal several indirect *f*CO₂-induced
48 effects on BPP and bacterial BV with potential consequences for oceanic carbon cycling, in
49 particular in a low nutrient and high *f*CO₂ future ocean.

50

51 **Key words**

52 Ocean acidification, CO₂ enrichment, Baltic Sea, KOSMOS mesocosm experiment, bacterial
53 production, phytoplankton, DOC accumulation



54 1 Introduction

55 Since the industrial revolution the oceans have absorbed ca. one half of the anthropogenic
56 carbon dioxide (CO₂), thereby shifting carbonate chemistry equilibria and pH (Caldeira and
57 Wickett, 2003; Raven et al., 2005; Sabine et al., 2004). During the last decade, the Baltic Sea,
58 experienced a pronounced decrease in pH (~0.1 pH units between 1993 and 2012,
59 International Council for the Exploration of the Sea, 2014). This corresponds to a 30%
60 increase in the concentration of H⁺ during this period (IPCC, 2007) with potential
61 consequences for organism physiology (Fabry et al., 2008, Taylor et al., 2012). At the same
62 time, autotrophic organisms can be fertilized by an enhanced CO₂ availability increasing the
63 production of particulate (POM) and dissolved organic matter (DOM) (Egge, et al., 2009;
64 Hein and Sand-Jensen, 1997; Losh et al., 2012; Riebesell et al., 2007). However, most CO₂
65 enrichment experiments studying natural plankton assemblages under variable nutrient
66 conditions do not reveal a consistent response of primary production to elevated CO₂ (e.g.
67 Engel, et al., 2005; Hopkinson et al., 2010; Riebesell et al., 2007). Nevertheless, not only the
68 amount, but also the stoichiometric composition of algal DOM and POM can be affected by
69 changes in *f*CO₂. For example, Riebesell et al. (2007) or Maat et al. (2014) reported an
70 increased stoichiometric drawdown of carbon (C) to nitrogen (N) at higher levels of *f*CO₂,
71 most likely as a result from C-overconsumption (Toggweiler, 1993). Since heterotrophic
72 bacteria greatly depend on phytoplankton derived organic carbon (e.g. Azam, 1998), they will
73 most likely respond to alterations in quantity and quality of phytoplankton derived DOM and
74 POM (e.g. Allgaier et al., 2008; Grossart et al., 2006a). Availability and competition for
75 nutrients, however, can substantially alter *f*CO₂-induced changes in activity and biomass of
76 phytoplankton and subsequently of heterotrophic bacteria. In nutrient-depleted or nutrient-
77 limited systems, bacteria can become restricted in their utilization of phytoplankton derived
78 organic matter, depending on the relative availability of inorganic nutrients (Hoikkala et al.,
79 2009; Lignell et al., 2008; Thingstad and Lignell, 1997). Consequently, a *f*CO₂ dependent
80 increase in inorganic C-availability for autotrophs may not stimulate heterotrophic activity.
81 This decoupling of heterotrophic from autotrophic processes has been termed as a
82 “counterintuitive carbon-to-nutrient coupling” (Thingstad et al., 2008). Consequently,
83 bioavailable dissolved organic carbon (DOC) and particulate organic carbon (POC) could
84 accumulate in nutrient limited oceanic surface waters with profound consequences for nutrient
85 cycling and the oceanic carbon pump (Cauwet et al., 2002; Mauriac et al., 2011; Søndergaard



86 et al., 2000; Thingstad et al., 1997). Various studies reported on limitation of bacterial growth
87 by inorganic nutrients in several parts of the Baltic Sea (e.g. Hoikkala et al., 2009; Kivi et al.,
88 1993; Kuparinen and Heinänen, 1993; Zweifel et al. 1993). Based on these results, we
89 intended to evaluate effects of enhanced $f\text{CO}_2$ on activity and biomass of free-living (FL) as
90 well as particle associated (PA) bacteria during a relatively low productive period of the year
91 with low levels of nutrients.

92

93 **2 Methods**

94 **2.1 Experimental setup, CO₂ manipulation and Sampling**

95 Nine floating, pelagic KOSMOS (Kiel Off-Shore Mesocosms for future Ocean Simulations;
96 Riebesell et al., 2013) were moored on 12th June 2012 (day -10 = t-10; 10 days before CO₂
97 manipulation) at 59°51.5'N, 23°15.5'E in the Baltic Sea at Tvärminne Storfjärden on the
98 south-west coast of Finland. Afterwards, the open mesocosm bags were rinsed and water fully
99 exchanged with the surrounding water masses for five days. Mesocosms were covered on the
100 top and bottom with a 3 mm net to exclude larger organisms. At t-5, sediment traps were
101 attached to the bottom at 17 m depth. Further, the submerged mesocosm bags were pulled up
102 1.5 m above the water surface, enclosing and separating ~55 m³ of water from the surrounding
103 Baltic Sea and covered by a photosynthetic active radiation (PAR) transparent roof to prevent
104 nutrient addition from birds and freshwater input from rain. Additionally, existing haloclines
105 were removed in each mesocosm as described in Paul et al. (2015), thereby creating a fully
106 homogeneous water body.

107 The experiment was conducted between 17th June (t-5) and 4th August (t43) 2012. CO₂
108 addition was performed stepwise on day t0 after sampling and the following three days to
109 minimize environmental stress on organisms until reaching the initial fugacity-levels of CO₂
110 ($f\text{CO}_2$). CO₂ addition was repeated at t15 in the upper mixed 7 m to compensate for
111 outgassing. Different $f\text{CO}_2$ treatments were achieved by equally distributing filtered (50 µm),
112 CO₂-saturated seawater into the treated mesocosms as described by Paul et al. (2015). Water
113 samples throughout the whole water column (0-17m) were collected from each mesocosm and
114 the surrounding seawater using depth-integrated water samplers (IWS, HYDRO-BIOS, Kiel).



115 Samples for activity measurements were directly subsampled from the IWS on the sampling
116 boat without headspace to maintain in-situ $f\text{CO}_2$ concentrations during incubation.

117 Unfortunately, three mesocosms were lost during the experiment due to welding faults and
118 thus unquantifiable water exchanges with the surrounding waters. Therefore, we only refer to
119 the six remaining mesocosms during this report, using the average $f\text{CO}_2$ from t1 to t43 to
120 characterize the different treatments as described in Paul et al. (2015): 365 μatm and 368
121 μatm (controls); 497 μatm , 821 μatm , 1007 μatm and 1231 μatm $f\text{CO}_2$, respectively. Detailed
122 descriptions on the study site, mesocosm deployment and system, performance of the
123 mesocosm facility throughout the experiment, CO_2 addition, carbonate chemistry, cleaning of
124 the mesocosm bags as well as sampling frequencies of single parameters can be obtained from
125 the experimental overview by Paul et al. (2015).

126 **2.2 Physical and chemical parameters**

127 Physical measurements (i.e. temperature and salinity) were performed using a CTC60M
128 memory probe (Sea and Sun Technology, Trappenkamp, Germany). For these parameters, the
129 depth-integrated mean values are presented. Full descriptions of sampling and analyses of
130 Chl *a*, particulate matter (particulate carbon (TPC), particulate organic nitrogen (PON), total
131 particulate phosphorus (TPP), biogenic silica (BSi)), dissolved organic matter (DOM
132 including dissolved organic carbon (DOC), dissolved organic nitrogen (DON), dissolved
133 organic phosphorus (DOP) as well as dissolved inorganic nutrients (phosphate (PO_4^{3-}),
134 nitrate (NO_3^-)) can be obtained from Paul et al. (2015) and in case of DOP measurements
135 from Nausch et al. (2015).

136 **2.3 Microbial standing stock**

137 Abundance of photoautotrophic cells (<20 μm) and free-living (FL) heterotrophic prokaryotes
138 (HP) were determined by flow cytometry (Crawford et al. 2016). In short, phytoplankton
139 were discriminated based on their chlorophyll red autofluorescence and/or phycoerythrin
140 orange autofluorescence (Marie et al., 1999). In combination with their side scatter signal and
141 size fractionation, the phytoplankton community could be divided into 6 clusters (Crawford et
142 al. 2016), varying in size from 1 to 8.8 μm average cell diameter. Three groups of
143 picoeukaryotic phytoplankton (Pico I-III), 1 picoprokaryotic photoautotroph (*Synechococcus*



144 spp.) and 2 nanoeukaryotic phytoplankton groups were detected. Biovolume (BV) estimations
145 were based on cell abundance and average cell diameters by assuming a spherical cell shape.
146 The BV sum of *Synechococcus* and Pico I-III is expressed as BV_{Pico} . The BV sum of Nano I
147 and II will be referred as BV_{Nano} . Abundances of FL HP were determined from 0.5 %
148 glutaraldehyde fixed samples after staining with a nucleic acid-specific dye (Crawford et al.
149 2016). Unicellular cyanobacteria (*Synechococcus* spp.) contributed at max 10% of the total
150 counts and, therefore, we use the term heterotrophic prokaryotes (HP). Two groups were
151 identified based on their low (LDNA) and high (HDNA) fluorescence.

152 Particle-associated (PA) HP were enumerated by epifluorescence-microscopy on a Leica
153 Leitz DMRB fluorescence microscope with UV- and blue light excitation filters (Leica
154 Microsystems, Wetzlar, Germany). Fresh samples were gently mixed to prevent particle
155 settling and a subsample of 15 mL was filtered on a 0.1-% Irgalan Black coloured 5.0 μm
156 polycarbonate-filter (Whatman, Maidstone, UK) (Hobbie et al., 1977). Thereafter, filters were
157 fixed with glutaraldehyde (Carl Roth, Karlsruhe, Germany, final conc. 2 %) and stained for 15
158 min with 4'-diamidino-2-phenylindole (DAPI, final conc. 1 $\mu\text{g mL}^{-1}$) (Porter and Feig, 1980)
159 directly on the filtration device and rinsed twice with sterile filtered habitat water before air-
160 drying and embedding in Citifluor AF1 (Citifluor Ltd, London, UK) on a microscopic slide
161 (Rieck et al. 2015). Due to mainly small, equally distributed particles on the filters throughout
162 the experiment, 15 random unique squares were counted with a magnification of 1000x. Total
163 number of PA HP was enumerated by subtracting autofluorescent cells from DAPI-stained
164 cells.

165 BV was calculated separately for FL and PA HP. For FL HP, we used an average cell volume
166 of 0.06 μm^3 reported by Hagström et al. (1979). BV of PA HP were calculated from
167 measurements of 1600 cells from 3 different mesocosms (346 μatm , 868 μatm , 1333 μatm) as
168 well as different time points throughout the experiment (t_0 , t_{20} , t_{39}) according to Massana et
169 al. (1997). The resulting average BV of 0.16 μm^3 per cell was further used to calculate BV of
170 PA HP from cell abundances. The BV-sum of both size fractions is expressed as total BV of
171 HP (BV_{HP}). Thereby, cell-numbers of PA HP were interpolated with R (R Core Team, 2014),
172 using splines, to calculate daily abundances. Further, we use the term “HP” and
173 “heterotrophic bacteria” synonymously, since heterotrophic bacteria account for the majority
174 of heterotrophic prokaryotes in surface waters (Karner et al., 2001; Kirchman et al. 2007).



175 Changes in Chl *a* and BV of heterotrophic bacteria are dependent on various factors, which
176 are not necessarily related to each other. Therefore, we have standardized BV_{HP} to total Chl *a*
177 known as a measurement for phytoplankton biomass (Falkowski and Kiefer, 1985). Thereby,
178 we express a ratio ($BV_{HP} : Chl a$), describing the distribution of heterotrophic bacterial BV and
179 phytoplankton biomass in relation to fCO_2 .

180 **2.4 Bacterial production and community respiration**

181 Rates of bacterial protein production (BPP) were determined by incorporation of ^{14}C -leucine
182 (^{14}C -Leu, Simon and Azam, 1989) according to Grossart et al. (2006a). Triplicates and a
183 formalin-killed control were incubated with ^{14}C -Leu (213 mCi $mmol^{-1}$; Hartmann Analytic
184 GmbH, Germany) at a final concentration of 165 nM, which ensured saturation of the uptake
185 systems of both FL and PA bacteria. Incubation was performed in the dark at *in situ*
186 temperature (between 7.8°C and 15.8°C) for 1.5 h. After fixation with 2% formalin, samples
187 were filtered onto 5.0 μm (PA bacteria) nitrocellulose filters (Sartorius, Germany) and
188 extracted with ice-cold 5% trichloroacetic acid (TCA) for 5 min. Thereafter, filters were
189 rinsed twice with ice-cold 5% TCA, once with ethanol (50% v/v), and dissolved in
190 ethylacetate for measurement by liquid scintillation counting (Wallac 1414, Perkin Elmer).
191 Afterwards, the collected filtrate was filtered on 0.2 μm (FL bacteria) nitrocellulose filters
192 (Sartorius, Germany) and processed in the same way as the 5.0 μm filters. Standard deviation
193 of triplicate measurements was usually <15%. The amount of incorporated ^{14}C -Leu was
194 converted into BPP by using an intracellular isotope dilution factor of 2. A conversion factor
195 of 0.86 was used to convert the produced protein into carbon (Simon and Azam, 1989). Cell-
196 specific BPP rates (csBPP) were calculated by dividing BPP-rates by abundances of PA and
197 FL HP.

198 Community respiration (CR) rates were calculated from oxygen consumption during an
199 incubation period of 48 hours at *in situ* temperature in the dark by assuming a respiratory
200 quotient of 1 (Berggren et al., 2012). Thereby oxygen concentrations were measured in
201 triplicate in 120 mL O_2 bottles without headspace, using a fiber optical dipping probe
202 (PreSens, Fibox 3), which was calibrated against anoxic and air saturated water. Further
203 descriptions are given by Spilling et al. (2015).



204 **2.5 Statistical analyses**

205 We used the nonparametric Spearman's rank correlation coefficient to measure statistical
206 dependence between variables. Significance is determined as $p < 0.05$. Statistical analyses and
207 visualisation were performed using R 3.1.2. (R Core Team, 2014) and R-package "ggplot2"
208 (Wickham, 2009).

209

210 **3 Results**

211 Paul et al. (2015) defined general phases of the experiment by physical characteristics of the
212 water column (temperature) as well as the first $f\text{CO}_2$ manipulation at t_0 (Phase 0 = t_{-5} to t_0 ,
213 Phase I = t_1 to t_{16} , Phase II = t_{17} to t_{30} , Phase III = t_{31} to t_{43}). These phases characterize
214 also changes in Chl *a* concentration and chemical bulk parameters. However, heterotrophic
215 bacteria differed in their response with a variable time delay. Consequently, we divided the
216 experiment into new phases based on changes in activity and BV of heterotrophic bacteria. To
217 provide clarity with respect to other publications of the same study, we termed the following
218 phases: **P1 = t_0 to t_8 , P2 = t_8 to t_{26} and P3 = t_{26} to t_{43}** . The time between closing of the
219 mesocosms and the first $f\text{CO}_2$ -manipulation was defined as Phase P0 = t_{-5} to t_0 . P1 describes
220 an initial phase without observed $f\text{CO}_2$ -related responses in BPP, csBPP or BV. During P2
221 several positive as well as negative $f\text{CO}_2$ -mediated effects on BPP, csBPP and BV were
222 observed, which could be related to the availability of phytoplankton derived organic carbon
223 and effects of bacterial mortality. The end of P2 is defined by reaching the BV maximum of
224 FL heterotrophic bacteria at t_{26} .

225 **3.1 Phytoplankton dynamics**

226 Concentration of Chl *a* increased after closing of the mesocosms until t_5 , followed by a
227 decline until the end of P1 (t_8) (Figure 1). During P0 and P1 no significant $f\text{CO}_2$ related
228 differences in total concentration of Chl *a* could be observed. During P2, concentrations of
229 Chl *a* increased again, driven by increasing BV of nanophotoautotrophs (BV_{Nano}) until
230 reaching the respective BV maximum of nanophotoautotrophs as well as Chl *a* at t_{16} - t_{17}
231 (Figure 1). Thereby, nanophotoautotrophs yielded significantly lower BV with increasing
232 $f\text{CO}_2$ between t_{13} - t_{17} ($r_s=0.68$, $p<<0.01$, $n=30$), which was reflected in lower concentrations
233 of Chl *a* in the 3 highest $f\text{CO}_2$ -treated mesocosms at the Chl *a* maximum at t_{16} . Thereafter,



234 both concentrations of Chl *a* and BV_{Nano} declined until t22-t28, respectively. During the
235 whole P2, Chl *a* was highly positively correlated to BV_{Nano} ($r_s=0.87$, $p<<0.01$, $n=123$). From
236 t22 until the end of the experiment, Chl *a* yielded overall low, but higher concentrations in the
237 3 highest $f\text{CO}_2$ -treated mesocosms ($r_s=0.71$, $p<<0.01$, $n=76$).

238 BV of picophotoautotrophs (BV_{Pico}) was positively correlated to overall Chl *a* development
239 during the initial phases P0 and P1 ($r_s=0.64$, $p<<0.1$, $n=66$), but showed a strong negative
240 correlation to Chl *a* during P2 and P3 ($r_s=-0.81$, $p<<0.1$, $n=162$). Especially after the
241 breakdown of Chl *a* at t16/t17, BV_{Pico} increased strongly towards the BV maximum at t24 and
242 remained constant until the end of the experiment (Figure 1). The increase was mainly driven
243 by BV of *Synechococcus* spp., which accounted for a generally high proportion of BV_{Pico}
244 ($31 \pm 2\%$ to $59 \pm 2\%$) during this study (Figure S1). All four groups of picophotoautotrophs
245 distinguished by flow cytometry, however, revealed positive or negative $f\text{CO}_2$ -related effects
246 on BV (Figure 2). During different periods the smallest sized photoautotroph Pico I ($\sim 1\ \mu\text{m}$)
247 as well as Pico II showed strong fertilization effects of $f\text{CO}_2$, whereas *Synechococcus* spp. and
248 Pico III were not and/or negatively affected by $f\text{CO}_2$.

249 **3.2 Bacterial production (BPP) and biovolume (BV)**

250 Heterotrophic bacterial BV was mainly made up by FL bacteria, as PA bacteria contributed to
251 only $2 \pm 0.7 - 10 \pm 0.7\%$ (mean $4.8 \pm 0.6\%$) of total bacterial BV. PA bacteria, however,
252 accounted for a substantial fraction of overall BPP ($27 \pm 1 - 59 \pm 7\%$, mean $39 \pm 4\%$). Both
253 bacterial size-fractions showed distinct dynamics in BV, BPP and csBPP during the course of
254 the experiment. Interestingly, we could not reveal any consistent and direct $f\text{CO}_2$ effect on
255 BPP, csBPP or BV of FL or PA heterotrophic bacteria. Nonetheless, we observed several
256 $f\text{CO}_2$ -related differences between the mesocosms in BPP of PA bacteria between t16 and t23
257 as well as BV, BPP and csBPP of FL bacteria within P2.

258 During the initial phases P0 and P1 changes in BPP and BV of both bacterial size-fractions
259 paralleled changes in Chl *a* and BV_{Pico}. Thereby, no significant differences or only weak
260 correlations in FL and PA bacterial BV as well as BPP and csBPP were observed with
261 changes in $f\text{CO}_2$ (Table 1). At t8, however, FL bacterial BPP and csBPP yielded 4-5 times
262 higher rates in the $f\text{CO}_2$ -treated mesocosms compared to both controls (Figure 3). These
263 higher FL BPP rates were well reflected in significantly higher BV of FL bacteria with



264 increasing $f\text{CO}_2$ from t10 to t13 ($r_s=0.72$; $p<<0.01$; $n=24$). Between t8-t13, FL bacterial BV
265 was positively correlated to BV_{Pico} ($r_s=0.52$, $p<<0.01$, $n=36$), but particularly to BV_{PicoI}
266 ($r_s=0.77$, $p<<0.01$, $n=36$). Surprisingly, after t13/t14, FL bacterial BV declined only in the
267 three highest $f\text{CO}_2$ -treated mesocosms until t18 (Figure 3). In parallel, BPP of both bacterial
268 size-fractions increased after the breakdown of Chl *a* at t16 and yielded significantly lower
269 rates at higher $f\text{CO}_2$ for PA bacteria ($r_s=-0.52$, $p<0.01$, $n=24$) as well as FL bacteria ($r_s=-0.51$,
270 $p=0.01$, $n=24$) between t16 and t26. Standardizing BPP rates to cell abundance, however,
271 revealed only significantly lower csBPP-rates at higher $f\text{CO}_2$ for FL bacteria during this
272 period ($r_s=-0.61$, $p<0.01$, $n=24$). Although we measured similar responses in BPP for PA and
273 FL bacteria between t16 and t26, BV of both size-fractions revealed contrasting dynamics
274 (Figure 3, Figure S2). PA bacterial BV declined with the decay of Chl *a*, whereas FL bacteria
275 increased strongly in BV, which was positively correlated to BV of picophotoautotrophs until
276 the end of P2. P3 was characterized by declining BPP rates and BV of heterotrophic bacteria.
277 FL or PA BPP, csBPP or BV were not or negatively correlated to Chl *a*, BV of
278 picophotoautotrophs or DOC during this period (Table 1).

279

280 4 Discussion

281 Although OA and its ecological consequences have received growing recognition during the
282 last decade (Riebesell and Gattuso, 2015), surprisingly little is known about the ecological
283 effects on heterotrophic bacterial biomass, production or microbial foodweb interactions at
284 nutrient depleted or nutrient limited conditions, since most of the experiments were carried
285 out during the productive phases of the year (e.g. phytoplankton blooms), under eutrophic
286 conditions (e.g. coastal areas), or even with nutrient additions (Allgaier et al., 2008; Brussaard
287 et al., 2013; Grossart et al., 2006a; Lindh et al., 2013; Riebesell, 2013). However, large parts
288 of the oceans are nutrient-limited or experience extended nutrient-limited periods during the
289 year (Moore et al., 2013). Thus, we conducted our experiment in July-August, when nutrients
290 and phytoplankton production were relatively low in the northeastern Baltic Sea (Hoikkala et
291 al., 2009; Lignell et al., 2008). During the study, low nitrogen availability limited overall
292 autotrophic production (Paul et al., 2015; Nausch et al., 2015). This resulted in a post spring
293 bloom phytoplankton community, dominated by picophytoplankton, which is known to
294 account for a large fraction of total phytoplankton biomass in oligotrophic, nutrient poor



295 systems (e.g. Agawin et al., 2000). Nevertheless, dynamics of Chl *a* revealed two minor
296 blooms of larger phytoplankton during the first half of the experiment. One developed directly
297 after the closing of the mesocosms, followed by a second one driven by nanophytoplankton
298 (Paul et al., 2015). Albeit, picophytoplankton accounted mostly > 50 % of Chl *a* during the
299 entire experiment (Paul et al., 2015). One reason might be, that picoplanktonic cells are
300 generally favoured compared to larger cells in terms of resource acquisition and subsequent
301 usage at low nutrient conditions due to their high volume to surface ratio as well as a small
302 boundary layer surrounding these cells (Moore et al., 2013; Raven, 1998). However, when
303 cell size is the major factor determining the access to dissolved nitrogen and phosphorous,
304 bacteria should be able to compete equally or better with picophytoplankton at low
305 concentrations (Drakare et al., 2003; Suttle et al., 1990). On the other hand, BV and
306 production of heterotrophic bacteria are highly dependent on quantity and quality of
307 phytoplankton-derived organic carbon and usually are tightly related to phytoplankton
308 development (Attermeyer et al., 2014; Attermeyer et al., 2015; Grossart et al., 2003; Grossart
309 et al., 2006b; Rösel and Grossart, 2012). Consequently, observed $f\text{CO}_2$ -induced effects on
310 phytoplankton abundance, phytoplankton losses due to grazing and viral lysis as well as $f\text{CO}_2$ -
311 related differences in phytoplankton composition altered the availability of phytoplankton-
312 derived organic matter for FL and PA heterotrophic bacteria (Crawford et al., 2016; Paul et
313 al., 2015). Subsequent, changes in BV and production of both size-fractions in relation to
314 differences in $f\text{CO}_2$ were observed. However, we could not reveal any consistent pattern of
315 $f\text{CO}_2$ -induced effects on the coupling of phytoplankton and bacteria. Changes in BV and
316 production of heterotrophic bacteria were rather indirectly related to different positive as well
317 as negative $f\text{CO}_2$ -correlated effects on the phytoplankton during relatively short periods.
318 These periods, however, comprised phases with high organic matter turnover (e.g. breakdown
319 of Chl *a* maximum). This notion emphasizes the importance to the oceanic carbon cycle,
320 especially during long periods of general low productivity. The last phase of the experiment
321 (P3), however, revealed also a decoupling of autotrophic production and heterotrophic
322 consumption, leading to relatively low, but still significantly higher accumulation of DOC at
323 enhanced $f\text{CO}_2$. Nonetheless, we observed additionally $f\text{CO}_2$ -mediated differences in FL
324 bacterial BV and cell-specific BPP rates, which could be related to effects of enhanced
325 bacterial grazing at higher $f\text{CO}_2$ (Crawford et al., 2016). Predicting effects on heterotrophic



326 bacteria in a future, acidified ocean might consequently depend on several complex trophic
327 interactions of heterotrophic bacteria within the pelagic food web.

328 **4.1 Bacteria-phytoplankton coupling at low nutrient concentrations**

329 Heterotrophic bacteria are important recyclers of autochthonously produced DOM in aquatic
330 systems and play an important role in nutrient regeneration in natural plankton assemblages
331 (Kirchman 1994, Brett et al., 1999). When phytoplankton is restricted in growth due to the
332 lack of mineral nutrients, often a strong commensalistic relationship between phytoplanktonic
333 DOM production and bacterioplanktonic DOM utilization has been observed (Azam et al.,
334 1983; Bratbak and Thingstad, 1985). Alterations in either growth conditions of phytoplankton
335 or DOM availability for heterotrophic bacterioplankton, but also losses of phyto- and
336 bacterioplankton due to grazing or viral lyses can influence the competition for nutrients and
337 DOM remineralization (Azam et al., 1983; Bratbak and Thingstad, 1985; Caron et al., 1988;
338 Sheik et al., 2014). The availability of DOM for heterotrophic bacteria may also change, when
339 they attach to living algae and organic particles. As a consequence, PA bacteria are often less
340 affected by nutrient limitation due to the generally higher nutrient availability at particle
341 surfaces (e.g. Grossart and Simon, 1993). In our study, this was reflected in the relatively high
342 csBPP rates of PA heterotrophic bacteria throughout the entire experiment. However, PA
343 heterotrophic bacteria contributed only a minor fraction (maximal 10 ± 0.7 %) to the overall
344 heterotrophic bacterial BV, which is usually reported for oligotrophic or mesotrophic
345 ecosystems (Lapoussière et al., 2010). Nevertheless, the substantial contribution of PA
346 heterotrophic bacteria to overall BPP emphasizes their importance, especially during such low
347 productive periods (e.g. Simon et al., 2002, Grossart, 2010). Generally, PA heterotrophic
348 bacteria are essential for the remineralization of nutrients from autotrophic biomass, which
349 would otherwise sink out from surface waters (Cho and Azam, 1988; Turley and Mackie,
350 1994). Leakage of hydrolysis products as well as attachment and detachment of bacteria to
351 and from particles stimulate production of the FL bacterial size fraction (Cho and Azam,
352 1988; Grossart et al., 2003, Smith et al., 1992) as well as equally-sized picophytoplankton,
353 which would be able to compete with bacteria in terms of nutrient-uptake. During the
354 breakdown of Chl *a* after t16/t17, both FL heterotrophic bacteria and picophotoautotrophs
355 benefitted from fresh, remineralized POM and their BV and production greatly increased



356 (Figure 3, Figure S2). The contrasting dynamics of PA heterotrophic bacteria might be a
357 result of particle losses via sinking (Turley and Mackie, 1994).

358 **4.1 $f\text{CO}_2$ -related effects on bacterial coupling to phytoplankton-derived** 359 **organic matter**

360 Several previous studies demonstrated that responses of heterotrophic bacteria due to changes
361 in $f\text{CO}_2$ were related to phytoplankton rather than being a direct effect of pH or CO_2 (e.g.
362 Allgaier et al., 2008, Grossart et al., 2006). Also during this study, BPP and BV of both
363 heterotrophic bacterial size-fractions were strongly linked to phytoplankton dynamics and
364 revealed several indirect responses to $f\text{CO}_2$, resulting from alterations in phytoplankton
365 community composition and biomass. One small picoeukaryote (Pico I) with cell-diameters of
366 $\sim 1 \mu\text{m}$ benefitted from the stepwise CO_2 addition, yielding significantly higher growth rates
367 and BV at higher $f\text{CO}_2$ after t3 (Crawfurd et al., 2016) (Figure 2). This is in line with a few
368 recent studies, indicating a positive effect of enhanced $f\text{CO}_2$ on the abundance of small
369 picoeukaryotic phytoplankton (Brussaard et al., 2013; Endo et al., 2013; Sala et al., 2015).
370 After t5, Pico I was controlled by grazing and viral lysis with highest reported viral lysis and
371 loss rates at t10 and t13, respectively (Crawfurd et al., 2016). Interestingly, viral lysis could
372 only be observed under high CO_2 conditions, but not at ambient CO_2 levels, which might be
373 related to higher Pico I productivity at increased $f\text{CO}_2$ (Crawfurd et al., 2016). Consequently,
374 at high $f\text{CO}_2$ biomass production of FL heterotrophic bacteria was fuelled by bioavailable
375 organic matter from viral lysis and grazing of algal cells (Brussaard et al., 1995; Brussaard et
376 al. 2005; Sheik et al., 2014). Thus, fertilization effects in photoautotrophic picoplankton
377 during CO_2 -addition and subsequent losses (Crawfurd et al., 2016) resulted indirectly in $f\text{CO}_2$ -
378 related differences in FL bacterial BV between t8 and t14 due to larger availability of
379 picophytoplankton-derived DOC.

380 In parallel a second phytoplankton-bloom developed, mainly driven by nanophytoplankton,
381 which yielded significantly lower BV at higher $f\text{CO}_2$ (Crawfurd et al., 2016). This was also
382 reflected in lower Chl *a* concentrations at highest $f\text{CO}_2$ (Paul et al., 2015). During breakdown
383 of Chl *a* after t16/t17, both BPP of FL and PA bacteria yielded significantly lower rates at
384 higher $f\text{CO}_2$, possibly due to the result of lower amounts of nanophytoplankton-derived
385 organic carbon. Nonetheless, differences in BV and csBPP dynamics of FL heterotrophic
386 bacteria between t14 and t26 could not be explained exclusively by the availability of



387 phytoplankton-derived organic carbon, but were rather caused by higher bacterial losses
388 mainly due to grazing at enhanced $f\text{CO}_2$ as reported by Crawford et al. (2016).

389 **4.2 Consequences of $f\text{CO}_2$ -related differences in bacterial mortality for** 390 **trophic relationships**

391 Not only heterotrophic bacterial activity but also mortality plays an important role in nutrient
392 regeneration in natural plankton assemblages (e.g. Caron 1994). Two major factors
393 determining bacterial mortality are viral lysis and grazing (e.g. Liu et al., 2010). The viral
394 shunt generates mainly bioavailable DOM and stimulates autotrophic and heterotrophic
395 microbes simultaneously. Advantages in competition for dissolved organic nutrients will
396 primarily benefit heterotrophic bacteria (e.g. Joint et al., 2002). In contrast, the consumption
397 of bacterial biomass by bacterivory may release phytoplankton from competition with bacteria
398 for limiting nutrients (e.g. Bratbak and Thingstad, 1985, Caron et al., 1990). Additionally,
399 carbon is directly transferred to higher trophic levels (Atkinson, 1996; Sherr et al., 1986;
400 Schnetzer and Caron, 2005). Both will certainly impact the tight phytoplankton-bacteria
401 coupling at low nutrient concentrations. However, possible effects of increased $f\text{CO}_2$ on the
402 impact of bacterial grazing for trophic interactions are so far largely unknown. Only a few
403 studies have reported on bacterial grazing in ocean acidification research under different
404 nutrient conditions and indicated both no effects as well as effects of $f\text{CO}_2$ (e.g. Brussaard et
405 al., 2013; Rose et al., 2009; Suffrian et al., 2008).

406 During our study FL heterotrophic bacterial BV surprisingly dropped only in the highest
407 $f\text{CO}_2$ -treated mesocosms after t13/t14 and stayed low until t22. In particular, the delay of FL
408 bacterial BV increase after the Chl *a* break-down at t16/t17 was rather long, since
409 heterotrophic bacteria usually react on much shorter time scales to alterations in
410 phytoplankton-derived organic matter (e.g. Azam et al., 1993). Crawford et al. (2016),
411 however, reported significantly higher bacterial grazing at enhanced $f\text{CO}_2$ from grazing assays
412 at t15. Consequently, higher availability of DOM after the decay of the phytoplankton bloom
413 did stimulate BPP, but this biomass production was directly channelled to a larger proportion
414 by grazing to higher trophic levels at enhanced $f\text{CO}_2$ (Atkinson, 1996; Schnetzer and Caron,
415 2005; Sherr et al., 1986). Nevertheless, we also may add viral lysis here as a possibility for a
416 higher bacterial mortality. Indeed, viral abundance was higher at enhanced $f\text{CO}_2$ but increased
417 already after t8 and remained on a constant level until t22 (Crawford et al., 2016). Although it



418 is unlikely that viral lysis caused the observed $f\text{CO}_2$ -related differences in bacterial BV
419 dynamics between t13/t14 and t26, it still might have added to some of the $f\text{CO}_2$ -related
420 effects during this period.

421 In addition, Crawford et al. (2016) reported following flow cytometry analysis an
422 accompanying drop of HDNA, but not LDNA bacteria between t13/t14 and t19, which altered
423 finally the proportion of HDNA:LDNA bacteria in relation to $f\text{CO}_2$ between t14 and t26.
424 Differentiation of LDNA and HDNA bacteria according to the cell's nucleic acid content can
425 indicate differences in cell size (Gasol and del Giorgio, 2000), but is more likely a measure
426 for the cell's activity (Gasol and del Giorgio, 2000; Lebaron et al., 2001; Schapira et al.,
427 2009). Although we cannot draw any conclusion, if cell size or cell-activity was finally the
428 determining factor, preferential grazing on HDNA heterotrophic bacteria seems likely (Gasol
429 et al., 1999, Hahn and Höfle, 2001; Vaqué, 2001). This resulted, however, in a higher
430 contribution of LDNA and possibly smaller as well as less active cells to the heterotrophic
431 bacterial population. At higher $f\text{CO}_2$ subsequent FL cell-specific BPP rates were reduced and
432 BPP maxima more delayed in time between t16 and t26.

433 Unfortunately, we are not able to relate our results to any possible group of grazing
434 organisms. Nevertheless, results from Flow Cytometry and counting of protozoa as well as
435 mesozooplankton indicated possible grazers (Bermúdez et al., 2016, Crawford et al., 2016,
436 Lischka et al., 2015). Bermúdez et al. (2016) reported highest biomass of protozoans around
437 t15. Biomass was thereby substantially made up by the heterotrophic choanoflagellate
438 *Calliacantha natans* (Bermúdez, pers. comm.). *Calliacantha natans* was demonstrated to feed
439 in a size-selective mode only on particles $< 1 \mu\text{m}$ in diameter (Marchant and Scott, 1993) and
440 thus could be a possible predator on heterotrophic bacteria. Additionally, Crawford et al.
441 (2016) distinguished one group of phototrophic picoeukaryotes by flow cytometry (Pico II),
442 which only increased in BV and thereby yielded significantly higher BV at higher $f\text{CO}_2$
443 during the period, when abundance of HDNA bacteria was reduced due to grazing. Although
444 we do not have any evidence for grazing of both particular groups of organisms, the type of
445 nutrition would have implications for trophic interactions. If the dominant grazers consisted of
446 mixotrophic organisms and would be able to fix carbon, they may have directly benefited
447 from increased CO_2 availability (Rose et al., 2009). Consequently, grazing on bacteria by
448 mixotrophs might have acted as a direct conduit for primary productivity supported by the use



449 of inorganic nutrients, which would otherwise be unavailable and bound in bacterial biomass
450 (Hartmann et al., 2012; Mitra et al. 2014; Sanders, 1991).

451 **4.3 Decoupling of $f\text{CO}_2$ -related effects on autotrophic production from** 452 **bacterial consumption during P3**

453 Exudation of carbon-rich substances by phytoplankton is one of the major sources of labile
454 DOM for heterotrophic bacteria (Larsson and Hagström, 1979). Exudation is highest under
455 nutrient-poor conditions, when nutrient limitation impedes phytoplankton growth, but not
456 photosynthetic carbon fixation (Fogg, 1983). Reported $f\text{CO}_2$ -related increases in primary-
457 production or in the consumption of inorganic carbon relative to nitrogen (e.g. Riebesell et al.,
458 1993, Riebesell et al., 2007) may potentially enhance exudation and subsequently alter
459 phytoplankton-bacteria interactions at higher $f\text{CO}_2$ (de Kluijver et al., 2010). During the last
460 phase of the experiment (P3) we indeed observed relatively low, but still significantly higher
461 DOC accumulation at enhanced $f\text{CO}_2$ (Figure 4). Although Spilling et al. (2016) could not
462 reveal any significant differences in primary production due to $f\text{CO}_2$, also pools of Chl *a* and
463 TPC as well as C:N_{POM} showed positive effects related to $f\text{CO}_2$ (Paul et al., 2015). However,
464 BPP and heterotrophic bacterial BV of both size-fractions did not reveal any similar $f\text{CO}_2$ -
465 related differences to DOC concentration or phytoplankton dynamics. This could lead to the
466 assumption, that heterotrophic bacteria were restricted in growth during P3. Similar findings
467 have been previously described by other studies, which reported on DOC-accumulation
468 caused by a limitation of DOM in surface waters (Cauwet et al., 2002; Larsen et al., 2015;
469 Mauriac et al., 2011; Thingstad et al., 1997, Thingstad et al., 2008). However, generally
470 strong increase in viral abundance and higher reported viral lysis of several phytoplankton
471 groups at higher $f\text{CO}_2$ would have also generated fresh bioavailable DOM during this period
472 (Crawford et al., 2016). Additionally, larger zooplankton increased strong in BV (Lischka et
473 al., 2015). Therefore an accumulation of DOC by escaping bacterial utilization seems likely,
474 since heterotrophic bacteria were possibly controlled by viral lysis and grazing. Nevertheless,
475 remineralized nutrients and carbon from the breakdown of the earlier phytoplankton blooms
476 were bound to a higher extend in autotrophic biomass at higher $f\text{CO}_2$ (Paul et al., 2015). This
477 is also reflected in a lower ratio of $\text{BV}_{\text{HP}} : \text{Chl}a$ with increasing $f\text{CO}_2$ (Figure 5). However,
478 during P3 $f\text{CO}_2$ -related differences did not impact sinking flux (Paul et al., 2015). This was
479 probably related to the domination of small-sized unicellular phytoplankton, which only



480 contributed indirectly via secondary processing of sinking material to the carbon export
481 (Richardson and Jackson, 2007, Paul et al., 2015). On the other hand, total CR rates were
482 significantly reduced at higher $f\text{CO}_2$ (Spilling et al., 2015) during P3. Interestingly, this
483 finding would suggest lower CR at higher DOC concentrations. However, CR was strongly
484 correlated to heterotrophic bacterial BV and thus reflected in the proportion of $\text{BV}_{\text{HP}} : \text{Chl } a$.
485 Consequently, the counterintuitive difference in CR during P3 is most likely a result of the
486 “heterotrophy” of the system, which was lower at higher $f\text{CO}_2$ (Figure 5).

487

488 5 Conclusion

489 Microbial processes can be affected either directly or indirectly via a cascade of effects
490 through the response of non-microbial groups or changes in water chemistry (Liu et al., 2010).
491 Our large-volume mesocosm approach allowed us to test for multiple $f\text{CO}_2$ -related effects on
492 heterotrophic bacterial activity and biovolume dynamics on a near-realistic ecosystem level
493 by including trophic interactions from microorganisms up to zooplankton. Thereby, we
494 addressed specifically a nutrient-depleted system, which is representative for large parts of the
495 oceans in terms of low nutrient concentrations and productivity (Moore et al., 2013). During
496 most time of the experiment, heterotrophic bacterial productivity was tightly coupled to the
497 availability of phytoplankton-derived organic matter and thus responded to $f\text{CO}_2$ -related
498 alterations in pico- and nanophytoplankton biovolume, albeit with contrasting results. So far,
499 this is the first ecosystem study, which cannot only report on positive, but also on
500 significantly negative effects of higher $f\text{CO}_2$ on bacterial production. During the experiment,
501 bacterial mortality from grazing and viral lysis had a strong impact on bacterial biovolume. In
502 particular, $f\text{CO}_2$ -induced effects on bacterial grazing and its impact on higher trophic levels
503 are still poorly understood and have been greatly neglected in ocean acidification research. In
504 our study, however, there was a period when autotrophic production was decoupled from
505 heterotrophic consumption, which resulted in a low, but significantly higher accumulation of
506 DOC, with potential consequences for carbon cycling in the upper ocean. Reasons and
507 consequences of these findings can unfortunately not be generalized, since we did not perform
508 specific bioassays to test for limiting nutrients. Thus, we highly encourage implementing such
509 bioassays during further experiments at low nutrient conditions. Our study reveals a number
510 of $f\text{CO}_2$ -induced effects, which led to responses in biovolume and productivity of



511 heterotrophic bacteria. Consequently, complex trophic interactions of heterotrophic bacteria in
512 the pelagic food web, which can only be successfully addressed in whole ecosystem studies,
513 seem to be the key for understanding and predicting $f\text{CO}_2$ -induced effects on aquatic food
514 webs and biogeochemistry in a future, acidified ocean.

515

516 **Acknowledgements**

517 We thank the KOSMOS team and all of the participants in the mesocosm campaign for
518 organisation, maintenance and support during the experiment. In particular, we would like to
519 thank Andrea Ludwig for coordinating the campaign logistics and assistance with CTD
520 operations and the diving team. Further we thank the Tvärminne Zoological Station for the
521 opportunity to carry out such a big mesocosm experiment at their research station and
522 technical support on site. Additionally we acknowledge the captain and crew of R/V *ALKOR*
523 for their work transporting, deploying (AL394) and recovering (AL397) the mesocosms. The
524 collaborative mesocosm campaign was funded by BMBF projects BIOACID II (FKZ
525 03F06550) and SOPRAN Phase II (FKZ 03F0611). CPDB was financially supported by the
526 Darwin project, the Royal Netherlands Institute for Sea Research (NIOZ), and the EU project
527 MESOAQUA (grant agreement number 228224).

528

529

530



531 **References**

- 532 Agawin, N.S.R., Duarte, C.M., Agusti, S.: Nutrient and temperature control of the
533 contribution of picoplankton to phytoplankton biomass and production, *Limnol. Oceanogr.*,
534 45 (3), 591-600, 2000.
- 535 Allgaier, M., Riebesell, U., Vogt, M., Thyrrhaug, R., Grossart, H.-P.: Coupling of
536 heterotrophic bacteria to phytoplankton bloom development at different $p\text{CO}_2$ levels: a
537 mesocosm study, *Biogeosciences*, 5, 1007-1022, 2008.
- 538 Atkinson, A.: Subantarctic copepods in an oceanic, low chlorophyll environment: ciliate
539 predation, food selectivity and impact on prey populations. *Mar Ecol Prog Ser.*, 130, 85–96,
540 1996.
- 541 Attermeyer, K., Hornick, T., Kayler, Z.E., Bahr, A., Zwirrmann, E., Grossart, H.-P., Premke,
542 K.: Enhanced bacterial decomposition with increasing addition of autochthonous to
543 allochthonous carbon without any effect on bacterial community composition.
544 *Biogeosciences*, 11 (6): 1479-1489, 2014.
- 545 Attermeyer, K., Tittel, J., Allgaier, M., Frindte, K., Wurzbacher, C.M., Hilt, S., Kamjunke, N.,
546 Grossart, H.-P.: Effects of light and autochthonous carbon additions on microbial turnover of
547 allochthonous organic carbon and community composition, *Microbial Ecology*, 69 (2): 361-
548 371, 2015.
- 549 Azam, F.: Microbial Control of Oceanic Carbon Flux: The Plot Thickens, *Science*, 280
550 (5364), 694-696, doi:10.1126/science.280.5364.694, 1998.
- 551 Azam, F., Fenchel, T., Field, J.G., Gray, J.S., Meyer-Reil, L.A., and Thingstad, F.: The
552 Ecological Role of Water-Column Microbes in the Sea, *Mar. Ecol. Prog. Ser.*, 10, 257-
553 263, 1983.
- 554 Azam, F., Smith, D.C., Steward, G.F., Hagström, Å.: Bacteria-Organic Matter Coupling and
555 Its Significance for Oceanic Carbon Cycling, *Microb. Ecol.*, 28, 167-179, 1993.
- 556 Berggren, M., Lapierre, J.-F., and del Giorgio, P. A.: Magnitude and regulation of
557 bacterioplankton respiratory quotient across fresh-water environmental gradients, *ISME*
558 *Journal*, 6, 984–993, 2012.



- 559 Bermúdez, J.R., Winder, M., Stuhr, A., Almén, A.K., Engström-Öst, J., and Riebesell, U.:
560 Effect of ocean acidification on the structure and fatty acid composition of a natural plankton
561 community in the Baltic Sea, Biogeosciences Discuss., doi:10.5194/bg-2015-669, 2016.
- 562 Bratbak, G., Thingstad, T.F.: Phytoplankton-bacteria interactions: an apparent paradox?
563 Analysis of a model system with both competition and commensalism. Mar. Ecol. Prog. Ser.,
564 25, 23-30, 1985.
- 565 Brett, M.T., Lubnow, F.S., Villar-Argaiz, M., Müller-Solger, A., and Goldman, C.R.: Nutrient
566 control of bacterioplankton and phytoplankton dynamics, Aquatic Ecology, 33, 135-145,
567 1999.
- 568 Brussaard, C.P.C., Riegman, R., Noordeloos, A.A.M., Cadée, G.C., Witte, H., Kop, A.J.,
569 Nieuwland, G., Van Duyl, F.C, Bak, R.-P.M.: Effects of grazing, sedimentation and
570 phytoplankton cell lysis on the structure of a coastal pelagic food web, Mar. Ecol. Prog. Ser.,
571 123, 259-271, 1995.
- 572 Brussaard, C.P.D., Kuipers, B., Veldhuis, M.J.W.: A mesocosm study of *Phaeocystis globosa*
573 population dynamics: I. Regulatory role of viruses in bloom, Harmful Algae, 4, 859–874,
574 2005.
- 575 Brussaard, C.P.C., Noordeloos, A.A.M., Witte, H., Collenteur, M.C.J., Schulz, K.G., Ludwig,
576 A., Riebesell, U.: Arctic microbial community dynamics influenced by elevated CO₂ levels,
577 Biogeosciences, 10, 719-731, 2013.
- 578 Caldeira, K. and Wickett, M.E.: Anthropogenic carbon and ocean pH, Nature, 425, 365, 2003.
- 579 Caron, D.A.: Inorganic Nutrients, Bacteria, and the Microbial Loop, Microb. Ecol., 28, 295-
580 298, 1994.
- 581 Caron DA, Goldman JC: Protozoan nutrient regeneration. In: Capriulo GM [ed] Ecology of
582 marine protozoa. Oxford University Press, New York, p 283–306, 1990.
- 583 Cauwet, G., Déliat, G., Krastev, A., Shtereva, G., Becqueevort, S., Lancelot, C., Momzikoff,
584 A., Saliot, A., Cociasu, A., Popa, L.: Seasonal DOC accumulation in the Black Sea: a regional
585 explanation for a general mechanism, Marine Chemistry, 79, 193-205, 2002.
- 586 Cho, B.C., Azam, F.: Major role of bacteria in biogeochemical fluxes in the ocean's interior,
587 Nature, 332, 441-443, 1988.



- 588 Crawford, K.J., Riebesell, U., and Brussaard, C.P.D.: Shifts in the microbial community in the
589 Baltic Sea with increasing CO₂, *Biogeosciences Discuss.*, doi:10.5194/bg-2015-606, 2016
- 590 de Kljijver, A., Soetaert, K., Schultz, K.-G., Riebesell, U., Bellerby, R.G.J., and Middelburg,
591 J.J.: Phytoplankton-bacteria coupling under elevated CO₂ levels: a stable isotope labelling
592 study, *Biogeosciences*, 7, 3783-3793, 2010.
- 593 Drakare, S., Blomqvist, P., Bergström, A.-K. and Jansson, M.: Relationships between
594 picophytoplankton and environmental variables in lakes along a gradient of water colour and
595 nutrient content, *Freshwater Biology*, 48, 729-740, 2003.
- 596 Egge, J.K., Thingstad, T.F., Larsen, A., Engel, A., Wohlers, J., Bellerby, R.G.J., Riebesell, U.:
597 Primary production during nutrient-induced blooms at elevated CO₂ concentrations,
598 *Biogeosciences*, 6, 877-885, 2009
- 599 Endo, H., Yoshimura, T., Kataoka, T., Suzuki, K.: Effects of CO₂ and iron availability on
600 phytoplankton and eubacterial community compositions in the northwest subarctic Pacific, *J.*
601 *Exp. Mar. Biol. Ecol.*, 439, 160-175, doi: 10.1016/j.jembe.2012.11.003, 2013.
- 602 Engel, A., Zondervan, I., Aerts, K., Beaufort, L., Benthien, A., Chou, L., Delille, B., Gattuso,
603 J.-P., Harlay, J., and Heemann, C.: Testing the direct effect of CO₂ concentration on a bloom
604 of the coccolithophorid *Emiliana huxleyi* in mesocosm experiments, *Limnol. Oceanogr.*, 50,
605 493–507, doi:10.4319/lo.2005.50.2.0493, 2005.
- 606 Fabry, V.J., Seibel, B.A., Feely, R.A., Orr, J.C.: Impacts of ocean acidification on marine
607 fauna and ecosystem processes, *ICES Journal of Marine Science*, 65, 414-432, 2008
- 608 Falkowski, P., Kiefer, D.A.: Chlorophyll *a* fluorescence in phytoplankton: relationship to
609 photosynthesis and biomass, *J. Plankton Res.*, 7 (5), 715-731, 1985, doi:
610 10.1093/plankt/7.5.715, 1985.
- 611 Fogg, G.E.: The ecological significance of extracellular products of phytoplankton
612 photosynthesis, *Bot. Mar.*, 26, 3–14, 1983.
- 613 Gasol, J.M. and del Giorgio, P.A.: Using flow cytometry for counting natural planktonic
614 bacteria and understanding the structure of planktonic bacterial communities, *Sci. Mar.*, 64
615 (2), 197-224, 2000.



- 616 Gasol, J. M., Zweifel, U. L., Peters, F., Fuhrman, J. A., and Hagström, Å.: Significance of size
617 and nucleic acid content heterogeneity as measured by flow cytometry in natural planktonic
618 bacteria, *Appl. Environ. Microbiol.*, 65, 4475–4483, 1999.
- 619 Grossart, H.-P.: Ecological consequences of bacterioplankton lifestyles: changes in concepts
620 are needed. *Environ. Microbiol. Rep.*, 2, 706–714. doi: 10.1111/j.1758-2229.2010.00179.x,
621 2010.
- 622 Grossart, H.-P. and Simon, M.: Limnetic macroscopic organic aggregates (lake snow):
623 Occurrence, characteristics, and microbial dynamics in Lake Constance, *Limnol. Oceanogr.*,
624 38, 532-546, 1993.
- 625 Grossart H.-P., Hietanen S., Ploug H.: Microbial dynamics on diatom aggregates in Øresund,
626 Denmark. *Marine Ecology Progress Series*, 249: 69-78, 2003.
- 627 Grossart, H.-P., Allgaier, M., Passow, U., Riebesell, U.: Testing the effect of CO₂
628 concentration on the dynamics of marine heterotrophic bacterioplankton, *Limnology and*
629 *Oceanography*, 51, 1-11, 2006a
- 630 Grossart, H.-P., Czub, G., and Simon, M.: Specific interactions of planktonic algae and
631 bacteria: Implications for aggregation and organic matter cycling in the sea, *Environ.*
632 *Microbiol.*, 8, 1074–1084, 2006b.
- 633 Hagström, Å., Larsson, U., Hörstedt, P., Normark, S.: Frequency of Dividing Cells, a New
634 Approach to the Determination of Bacterial Growth Rates in Aquatic Environments, *Appl.*
635 *Environ. Microbiol.*, 37 (5), 805-812, 1979.
- 636 Hahn, M.W., Hofle, M.G.: Grazing of protozoa and its effect on populations of aquatic
637 bacteria, *FEMS Microbiology Ecology*, 35, 113-121-2001.
- 638 Hartmann, M., Grob, C., Tarran, G.A., Martin, A.P., Burkill, P.H., Scanlan, D.J. and Zubkov,
639 M.V.: Mixotrophic basis of Atlantic oligotrophic ecosystems, *Proc. Natl. Acad. Sci. U.S.A.*,
640 109 (15), 5756-5760, doi:10.1073/pnas.1118179109, 2012.
- 641 Hein, M. and Sand-Jensen, K.: CO₂ increases oceanic primary production, *Nature*, 388, 526-
642 527, doi:10.1038/41457, 1997.
- 643 Hobbie, J.E., Daley, R.J., Jasper, S.: Use of nuclepore filters for counting bacteria by
644 fluorescence microscopy, *Appl. Environ. Microbiol.*, 33, 1225-1228, 1977.



- 645 Hoikkala, L., Aarnos, H., Lignell, R.: Changes in Nutrient and Carbon Availability and
646 Temperature as Factors Controlling Bacterial Growth in the Northern Baltic Sea., *Estuaries
647 and Coasts*, 32, 720-733, doi:10.1007/s12237-009-9154-z, 2009.
- 648 Hopkinson, B. M., Xu, Y., Shi, D., McGinn, P. J., and Morel, F. M. M.: The effect of CO₂ on
649 the photosynthetic physiology of phytoplankton in the Gulf of Alaska, *Limnol. Oceanogr.*, 55,
650 2011–2024, doi:10.4319/lo.2010.55.5.2011, 2010.
- 651 Intergovernmental Panel on Climate Change (IPCC), *Climate Change 2007: The Scientific
652 Basis. Contribution of Working Group I to the Fourth Assessment Report of the
653 Intergovernmental Panel on Climate Change*, Solomon, S., Qin, D., Manning, M., Marquis,
654 M., Averyt, K., Tignor, M.M.B., Miller, H.L. (Eds.), Cambridge Univ. Press, New York,
655 2007.
- 656 International Council for the Exploration of the Sea: ICES Dataset on Ocean Hydrography,
657 ICES Oceanography Baltic Sea Monitoring Data, available at:
658 <http://ocean.ices.dk/helcom/Helcom.aspx?Mode=1>, last access: 7 August 2014
- 659 Joint, I., Henriksen, P., Fonnes, G.A., Bourne, D., Thingstad, T.F., Riemann, B.: Competition
660 for inorganic nutrients between phytoplankton and bacterioplankton in nutrient manipulated
661 mesocosms, *Aquat. Microb. Ecol.*, 29, 145-159, 2002.
- 662 Karner, M.B., DeLong, E.F. and Karl, D.M.: Archaeal dominance in the mesopelagic zone of
663 the Pacific Ocean, *Nature*, 409, 507-510, 2001.
- 664 Kirchman, D.L.: The Uptake of Inorganic Nutrients by Heterotrophic Bacteria, *Microb. Ecol.*,
665 28, 255-271, 1994
- 666 Kirchman, D.L., Elifantz, H., Dittel, A.I., Malmstrom, R.R. and Cottrell, M.T.: Standing
667 stocks and activity of Archaea and Bacteria in the western Arctic Ocean, *Limnol. Oceanogr.*,
668 52 (2), 495-507, 2007.
- 669 Kivi, K., Kaitala, S., Kuosa, H., Kuparinen, J., Leskinen, E., Lignell, R., Marcussen, B. and
670 Tamminen, T.: Nutrient limitation and grazing control of the Baltic plankton community
671 during annual succession. *Limnology and Oceanography* 38: 893–905, 1993.
- 672 Kuparinen, J. and Heinänen, A.: Inorganic Nutrient and Carbon Controlled Bacterioplankton
673 Growth in the Baltic Sea. *Estuarine, Coastal and Shelf Science* 37: 271–285, 1993.



- 674 Lapoussière, A., Michel, C., Starr, M., Gosselin, M., Poulin, M.: Role of free-living and
675 particle-attached bacteria in the recycling and export of organic material in the Hudson Bay
676 system, *Journal of Marine Systems*, 88, 434-445, 2011.
- 677 Larsen, A., Egge, J.K., Nejstgaard, J.C., Capua, I.D., Thyrrhaug, R., Bratbak, G., Thingstad,
678 T.F.: Contrasting response to nutrient manipulation in Arctic mesocosms are reproduced by
679 minimum microbial food web model, *Limnol. Oceanogr.*, 60, 360-374, 2015.
- 680 Larsson, U. and Hagström, Å.: Phytoplankton Exudate Release as an Energy-Source for the
681 Growth of Pelagic Bacteria, *Mar. Biol.*, 52, 199–206, 1979.
- 682 Lebaron, P., Servais, P., Agogue, H., Courties, C., and Joux, F.: Does the high nucleic acid
683 content of individual bacterial cells allow us to discriminate between active cells and inactive
684 cells in aquatic systems?, *Appl. Environ. Microbiol.*, 67, 1775–1782, 2001.
- 685 Lignell, R., Hoikkala, L., Lahtinen, T.: Effects of inorganic nutrients, glucose and solar
686 radiation treatments on bacterial growth and exploitation of dissolved organic carbon and
687 nitrogen in the northern Baltic Sea. *Aquat. Microb. Ecol.*, 51, 209–221, 2008.
- 688 Lindh, M.V., Riemann, L., Balter, F., Romero-Oliva, C., Salomon, P.S., Graneli, E. and
689 Pinhassi, J.: Consequences of increased temperature and acidification on bacterioplankton
690 community composition during a mesocosm spring bloom in the Baltic Sea, *Environmental*
691 *Microbiology Reports*, 5, 252-262, 2013.
- 692 Lischka, S., Bach, L.T., Schultz, K.-G., and Riebesell, U.: Micro- and mesozooplankton
693 community response to increasing CO₂ levels in the Baltic Sea: insights from a large-scale
694 mesocosm experiment, *Biogeosciences Discuss.*, 12, 20025-20070, doi:10.5194/bgd-12-20025-
695 2015, 2015.
- 696 Liu, J., Weinbauer, M.G., Maier, C., Dai, M., Gattuso, J.-P.: Effect of ocean acidification on
697 microbial diversity and on microbe-driven biogeochemistry and ecosystem functioning,
698 *Aquat. Microb. Ecol.*, doi:10.3354/ame01446, 2010.
- 699 Losh, J.L., Morel, F.M.M., Hopkinson, B.M.: Modest increase in the C:N ratio of N-limited
700 phytoplankton in the California Current in response to high CO₂, *Mar. Ecol.-Prog. Ser.*, 468,
701 31-42, doi:10.3354/meps09981, 2012.
- 702 Maat, D.S., Crawford, K.J., Timmermans, K.R. and Brussaard, C.P.D.: Elevated CO₂ and
703 Phosphate Limitation Favor *Micromonas pusilla* through Stimulated Growth and Reduced



- 704 Viral Impact, *Appl. Environ. Microbiol.*, 80 (10), 3119-3127, doi:10.1128/AEM.03639-13,
705 2014.
- 706 Marchant, H.J, Scott, F.J.: Uptake of sub-micrometre particles and dissolved organic material
707 by Antarctic choanoblagellates, *Mar. Ecol. Prog. Ser.*, 92, 59-64, 1993.
- 708 Marie, D., Brussaard, C.P.D., Thyraug, R., Bratbak, G., Vaultot, D.: Enumeration of marine
709 viruses in culture and natural samples by flow cytometry, *Appl. Environ. Microbiol.*, 65(1),
710 45-52, 1999.
- 711 Massana, R., Gasol, J.M., Bjørnsen, P.K., Blackburn, N., Hagström, Å, Hietanen, S., Hygum,
712 B.H., Kuparinen and Pedrós-Alió C.: Measurement of bacterial size via image analysis of
713 epifluorescence preparations: description of an inexpensive system and solutions to some of
714 the most common problems, *Sci. Mar.*, 61(3), 397-407, 1997.
- 715 Mauriac, R., Moutin, T., Baklouti, M.: Accumulation of DOC in Low Phosphate Low
716 Chlorophyll (LPLC) area: is it related to higher production under high N:P ratio?,
717 *Biogeosciences*, 8, 933-950, 2011.
- 718 Mitra, A., Flynn, K.J., Burkholder, J.M., Berge, T., Calbet, A., Raven, J.A., Granéli, E.,
719 Gilbert, P.M., Hansen, P.J., Stoecker, D.K., Thingstad, F., Tillmann, U., Våge, S., Wilken, S.,
720 and Zubkov, M.V.: The role of mixotrophic protists in the biological carbon pump,
721 *Biogeosciences*, 11, 995-1005, doi:10.5194/bg-11-995-2014, 2014.
- 722 Moore, C.M., Mills, M.M., Arrigo, K.R., Berman-Frank, I, Bopp, L., Boyd, P.W., Galbraith,
723 E.D., Geider, R.J., Guieu, C., Jaccard, S.L., Jickells, T.D., La Roche, J., Lenton, T.M.,
724 Mahowald, N.M., Marañón, E., Marinov, I, Moore, J.K., Nakatsuka, T., Oschlies, A., Sito,
725 M.A., Thingstad, T.F., Tsuda, A. and Ulloa, O.: Processes and patterns of oceanic nutrient
726 limitation, *Nature Geoscience*, 6(9), 701-710, doi:10.1038/NGEO1765, 2013
- 727 Nausch, M., Bach, L., Czerny, J., Godstein, J., Grossart, H.-P., Hellemann, D., Hornick, T.,
728 Achterberg, E., Schulz, K.G and Riebesell, U.: Effects of CO₂ perturbation on phosphorus
729 pool sizes and uptake in a mesocosm experiment during a low productive summer season in
730 the northern Baltic Sea, *Biogeosciences Discuss.*, 12, 17543-17593, doi:10.5194/bgd-12-
731 17543-2015, 2015.
- 732 Paul, A.J., Bach, L.T., Schulz, K.-G., Boxhammer, T., Czerny, J., Achterberg, E.P.,
733 Hellemann, D., Trense, Y., Nausch, M., Sswat, M., Riebesell, U.: Effect of elevated CO₂ on



- 734 organic matter pools and fluxes in a summer Baltic Sea plankton community, *Biogeosciences*,
735 12, 1-23, doi:10.5194/bg-12-1-2015, 2015.
- 736 Porter, K.G., Feig, Y.S.: Dapi for identifying and counting aquatic microflora, *Limnol.*
737 *Oceanogr.*, 25, 943-948, 1980
- 738 Raven, J.A.: The twelfth Tansley Lecture. Small is beautiful: the picophytoplankton,
739 *Functional Ecology*, 12, 503-513, 1998.
- 740 Raven, J., Caldeira, K., Elderfield, H., Hoegh-Guldberg, O., Liss, P., Riebesell, U.,
741 Shepherd, J., Turley, C., Watson, A.: *Ocean Acidification due to Increasing Atmospheric*
742 *Carbon Dioxide*, The Royal Society, London, UK, 2005.
- 743 R Core Team (2014). *R: A language and environment for statistical computing*. R Foundation
744 for Statistical Computing, Vienna, Austria. URL <http://www.R-project.org/>, 2014.
- 745 Richardson, T.L. and Jackson, G.A.: Small phytoplankton and carbon export from the surface
746 ocean, *Science*, 315, 838–840, doi:10.1126/science.1133471, 2007.
- 747 Riebesell, U., Wolfgladrow, D. A., and Smetacek, V.: Carbondioxide limitation of marine-
748 phytoplankton growth-rates, *Nature*, 361, 249–251, 1993.
- 749 Riebesell, U., Gattuso, J.-P.: Lessons learned from ocean acidification research. Reflection on
750 the rapidly growing field of ocean acidification research highlights priorities for future
751 research on the changing ocean. *Nature Climate Change*, 5, 12-14, doi:10.1038/nclimate2456,
752 2015.
- 753 Riebesell, U., Schulz, K.G., Bellerby, R.G.J., Botros, M., Fritsche, P., Meyerhöfer, M., Neill,
754 C., Nondal, G., Oschlies, A., Wohlers and J., Zöllner, E.: Enhanced biological carbon
755 consumption in a high CO₂ ocean, *Nature Letters*, 450 (22), 545-548,
756 doi:10.1038/nature06267, 2007.
- 757 Riebesell, U., Gattuso, J.-P., Thingstad, T.F. and Middelburg, J.J.: Arctic ocean acidification:
758 pelagic ecosystem and biogeochemical responses during a mesocosm study, *Biogeosciences*,
759 10, 5619-5626, doi:10.5194/bg-10-5619-2013, 2013.
- 760 Rieck, A., Herlemann, D.P.R., Jürgens, K. and Grossart, H.-P.: Particle-Associated Differ
761 from Free-Living Bacteria in Surface Waters of the Baltic Sea, *Front. Microbiol.*, 6 (1297),
762 doi:10.3389/fmicb.2015.01297, 2015.



- 763 Rose, J.M., Feng, Y., Gobler, C.J., Gutierrez, R., Hare, C.E., Leblanc, K., Hutchins, D.A.:
764 Effects of increased pCO₂ and temperature on the North Atlantic spring bloom. II.
765 Microzooplankton abundance and grazing, *Mar. Ecol. Prog. Ser.*, 388, 27-40, 2009.
- 766 Rösel, S., Grossart, H.-P.: Contrasting dynamics in activity and community composition of
767 free-living and particle-associated bacteria in spring, *Aquatic Microbial Ecology*, 66 (1), 169-
768 181, 2012.
- 769 Sabine, C. L., Feely, R. A., Gruber, N., Key, R. M., Lee, K., Bullister, J. L., Wanninkhof,
770 R., Wong, C. S., Wallace, D.W., Tilbrook, B., Millero, F. J., Peng, T. H., Kozyr, A., Ono, T.,
771 and Rios, A. F.: The oceanic sink for anthropogenic CO₂, *Science*, 305, 367–371, 2004.
- 772 Sala, M.M., Aparicio, F.L., Balagué, V., Boras, A., Borrull, E., Cardelús, C., Cros, L., Gomes,
773 A., López-Sanz, A., Malits, A., Martínez, R.A., Mestre, M., Movilla, J., Sarmiento, H.,
774 Vázquez-Dominguez, E., Vaqué, D., Pinhassi, J., Calbet, A., Calvo, E., Gasol, J.M., Pelejero,
775 C., Marrasé, C.: Contrasting effects of ocean acidification on the microbial food web under
776 different trophic conditions, *ICES Journal of Marine Science*, doi:10.1093/icesjms/fsv130,
777 2015.
- 778 Sanders, R.W.: Mixotrophic Protists in Marine and Freshwater Ecosystems, *J. Protozool.*, 38
779 (1), 76-81, 1991.
- 780 Schapira, M., Pollet, T., Mitchell, J.G. and Seuront, L.: Respiration rates in marine
781 heterotrophic bacteria relate to the cytometric characteristics of bacterioplankton
782 communities, *Journal of the Marine Biological Association of the United Kingdom*, 89 (6),
783 1161-1169, doi:http://dx.doi.org/10.1017/S0025315409000617, 2009.
- 784 Schnetzer, A., Caron, D.A.: Copepod grazing impact on the trophic structure of the microbial
785 assemblage of the San Pedro Channel, California; *J. Plankton Res.*, 27, 959–972, 2005.
- 786 Sheik, A.R., Brussaard, C.P.D., Lavik, G., Lam, P., Musat, N., Krupke, A., Littmann, S.,
787 Strous, M. and Kuypers M.M.M.: Responses of the coastal bacterial community to viral
788 infection of the algae *Phaeocystis globosa*, *The ISME Journal*, 8, 212-225, doi:
789 10.1038/ismej.2013.135, 2014.
- 790 Sherr, E.B., Sherr, B.F., Paffenhöfer, G.A.: Phagotrophic protozoa as food for metazoans: a
791 ‘missing’ trophic link in marine food webs, *Mar. Microb. Food Webs*, 1, 61–80, 1986.



- 792 Simon, M., Azam, F.: Protein content and protein synthesis rates of planktonic marine
793 bacteria, *Marine Ecology Progress Series*, 51, 201-213, 1989.
- 794 Simon, M., Grossart, H.-P., Schweitzer, B., and Ploug, H.: Microbial ecology of organic
795 aggregates in aquatic ecosystems. *Aquat. Microb. Ecol.*, 28, 175–211,
796 doi:10.3354/ame028175, 2002.
- 797 Smith, D.C., Simon, M., Alldredge, A.L., Azam, F.: Intense hydrolytic enzyme activity on
798 marine aggregates and implications for rapid particle dissolution, *Nature*, 359, 139-142, 1992.
- 799 Søndergaard, M., Williams, P. le B., Cauwet, G., Riemann, B., Rabinson, C., Terzic, S.,
800 Woodward, E.M.S., Worm, J.: Net accumulation and flux of dissolved organic carbon and
801 dissolved organic nitrogen in marine plankton communities, *Limnol. Oceanogr.*, 45(5), 1097-
802 1111, 2000.
- 803 Spilling, K., Paul, A.J., Virkkala, N., Hastings, T., Lischka, S., Stuhr, A., Bermudéz, R.,
804 Czerny, J., Boxhammer, T., Schulz, K.G., Ludwig, A., and Riebesell, U.: Ocean acidification
805 decreases plankton respiration: evidence from a mesocosm experiment, *Biogeosciences*
806 Discuss., doi:10.5194/bg-2015-608, 2016.
- 807 Suffrian, K., Simonelli, P., Nejstgaard, J.C., Putzeys, S., Carotenuto, Y., and Antia, A.N.:
808 Microzooplankton grazing and phytoplankton growth in marine mesocosms with increased
809 CO₂ levels. *Biogeosciences*, 5, 1145-1156, 2008.
- 810 Suttle, C.A., Fuhrman, J.A., Capone, D.G.: Rapid ammonium cycling and concentration-
811 dependent partitioning of ammonium and phosphate: Implications for carbon transfer in
812 planktonic communities, *Limnol. Oceanogr.*, 35 (2), 424-433, 1990
- 813 Taylor, A.R., Brownlee, C., Wheeler, G.L.: Proton channels in algae: reasons to be excited,
814 *Trends in Plant Sciences*, 17(11), 675-684, doi:10.1016/j.tplants.2012.06.009, 2012
- 815 Thingstad, T.F., and R. Lignell, R.: Theoretical models for the control of bacterial growth
816 rate, abundance, diversity and carbon demand. *Aquat. Microb. Ecol.*, 13, 19–27, 1997.
- 817 Thingstad, T.F., Hagström, Å., Rassoulzadegan, F.: Accumulation of degradable DOC in
818 surface waters: It is caused by a malfunctioning microbial loop?, *Limnol. Oceanogr.*, 42(2),
819 398-404, 1997.
- 820 Thingstad, T.F., Bellerby, R.G.J., Bratbak, G., Borsheim, K.Y., Egge, J.K., Heldal, M.,
821 Larsen, A., Neill, C., Nejstgaard, J., Norland, S., Sandaa, R.-A., Skjoldal, E.F., Tanaka, T.,



- 822 Thyrrhaug, R., Töpper, B.: Counterintuitive carbon-to-nutrient coupling in an Arctic pelagic
823 ecosystem, *Nature Letters*, 455, 387-391, doi:10.1038/nature07235, 2008.
- 824 Toggweiler, J.R.: Carbon overconsumption, *Nature*, 363, 210-211, 1993.
- 825 Turley, C.M., Mackie, P.J.: Biogeochemical significance of attached and free-living bacteria
826 and the flux of particles in the NE Atlantic Ocean., *Mar. Ecol. Prog. Ser.*, 115; 191-203, 1994.
- 827 Vaqué, D., Casamayor, E. O., and Gasol, J. M.: Dynamics of whole community bacterial
828 production and grazing losses in seawater incubations as related to the changes in the
829 proportions of bacteria with different DNA content, *Aquat. Microb. Ecol.*, 25, 163–177, 2001.
- 830 Wickham, H.: *ggplot2: elegant graphics for data analysis*. Springer New York, 2009.
- 831 Zweifel, U.L., Norrman, B., and Hagström, Å.: Consumption of dissolved organic carbon by
832 marine bacteria and demands for inorganic nutrients. *Marine Ecology Progress Series* 101,
833 23–32, 1993.



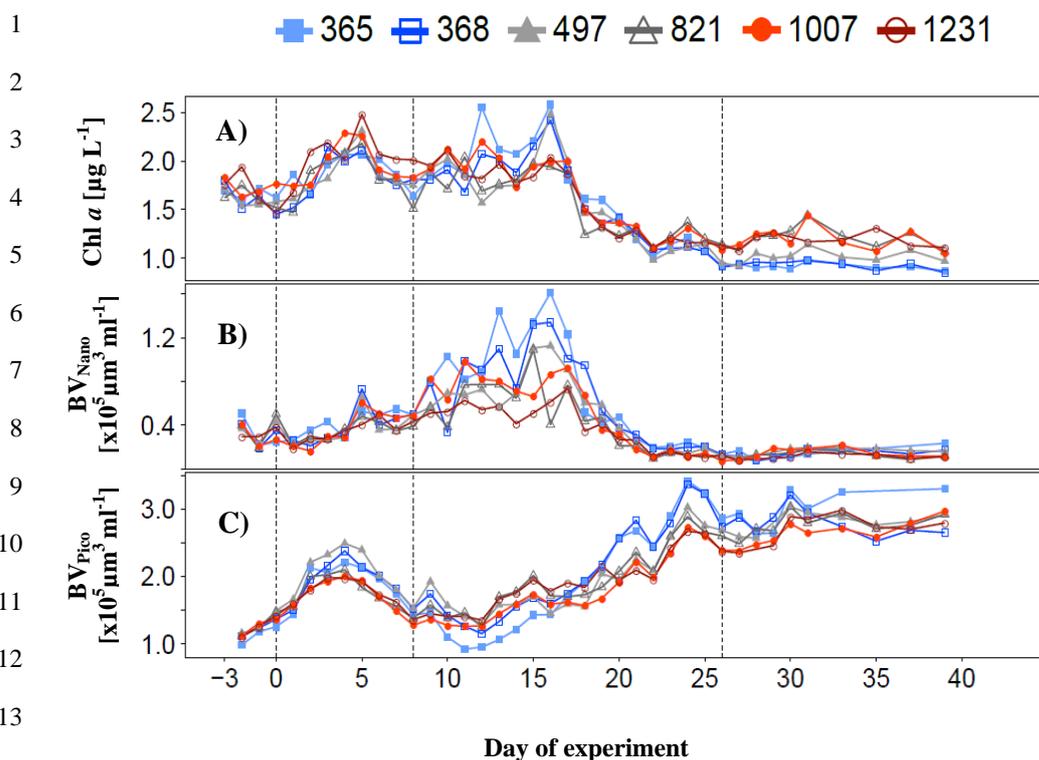
1 Table 1: Spearman's rank correlation (Spearman's rank correlation coefficient r_s ; p-value; n)
 2 of heterotrophic prokaryotic biovolume (BV_{HP}), bacterial protein production (BPP) and cell-
 3 specific BPP of size-fractions I) 0.2-5.0 μm (free-living; FL) and II) $>5.0 \mu\text{m}$ (particle-
 4 associated; PA) with $f\text{CO}_2$, dissolved organic carbon (DOC), community respiration (CR),
 5 chlorophyll *a* (Chl *a*) and total as well as group-specific biovolumes of pico- and
 6 nanophotoautotrophs (*Synechococcus* spp, Pico I-III, Nano I-II) during the different phases of
 7 the experiment. (n.s.- not significant)

8

	I) FL size fraction			II) PA size fraction		
	BV_{HP}	BPP	csBPP	BV_{HP}	BPP	csBPP
$f\text{CO}_2$	P0: -	P0: -	P0: -	P0: -	P0: -	P0: -
	P1: 0.36; 0.01; 48	P1: 0.5; 0.01; 24	P1: 0.55; <<0.01; 24	P1: n.s.	P1: n.s.	P1: 0.41; 0.05; 24
	P2: n.s.	P2: n.s.	P2: n.s.	P2: n.s.	P2: n.s.	P2: n.s.
	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.
DOC	P0: -	P0: -	P0: -	P0: -	P0: -	P0: -
	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.
	P2: n.s.	P2: 0.65; <0.01; 17	P2: n.s.	P2: n.s.	P2: 0.49; 0.05; 17	P2: n.s.
	P3: n.s.	P3: -0.35; 0.02; 44	P3: -0.35; 0.03; 38	P3: n.s.	P3: n.s.	P3: n.s.
CR	P0: -0.71; <0.01; 12	P0: n.s.	P0: n.s.	P0: -0.62; 0.03; 12.	P0: n.s.	P0: n.s.
	P1: 0.58; <<0.01; 42	P1: n.s.	P1: n.s.	P1: 0.5; 0.03; 18	P1: n.s.	P1: n.s.
	P2: 0.64; <<0.01; 106	P2: 0.72; <<0.01; 36	P2: 0.51; <0.01; 36	P2: 0.5; <0.01; 36	P2: 0.71; <<0.01; 36	P2: n.s.
	P3: 0.59; <<0.01; 36	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.
Chl <i>a</i>	P0: n.s.	P0: -0.59; 0.04; 12	P0: -0.89; 0.02; 6	P0: -0.65; 0.02; 12	P0: n.s.	P0: n.s.
	P1: 0.77; <<0.001; 48	P1: 0.48; 0.02; 24	P1: n.s.	P1: 0.39; 0.05; 24	P1: 0.51; 0.01; 24	P1: n.s.
	P2: -0.77; <<0.001; 112	P2: -0.41; <0.01; 41	P2: n.s.	P2: n.s.	P2: -0.49; <0.01; 41	P2: -0.41; 0.01; 41
	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: -0.31; 0.05; 41
BV_{Nano}	P0: n.s.	P0: n.s.	P0: n.s.	P0: n.s.	P0: 0.83; 0.04; 6	P0: n.s.
	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.
	P2: -0.75; <<0.01; 112	P2: -0.35; 0.02; 42	P2: n.s.	P2: n.s.	P2: -0.44; <0.01; 42	P2: 0.34; 0.03; 42
	P3: -0.46; <<0.01; 51	n.s.	P3: 0.35; 0.05; 33	P3: -0.32; 0.05; 39	P3: n.s.	P3: n.s.
BV_{Pico}	P0: 0.74; <0.01; 12	P0: n.s.	P0: n.s.	P0: n.s.	P0: n.s.	P0: n.s.
	P1: 0.79; <<0.01; 48	P1: 0.52; <0.01; 24	P1: n.s.	P1: 0.71; <<0.01; 24	P1: 0.58; <0.01; 24	P1: n.s.
	P2: 0.91; <<0.01; 112	P2: 0.65; <<0.01; 42	P2: n.s.	P2: 0.31; 0.04; 42	P2: 0.73; <<0.01; 42	P2: 0.37; 0.01; 42
	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.



BV_{Syn}	P0: 0.87; <<0.01; 12	P0: n.s.	P0: n.s.	P0: n.s.	P0: 0.83; 0.04; 6	P0: n.s.
	P1: 0.86; <<0.01; 48	P1: 0.5; 0.01; 24	P1: n.s.	P1: 0.64; <<0.01; 24	P1: 0.55; <0.01; 24	P1: n.s.
	P2: 0.89; <<0.01; 112	P2: 0.56; <<0.01; 42	P2: n.s.	P2: n.s.	P2: 0.55; <<0.01; 42	P2: 0.37; 0.01; 42
	P3: n.s.	P3: -0.44; <0.01; 38	P3: -0.47; <0.01; 33	P3: n.s.	P3: -0.5; <0.01; 38	P3: n.s.
BV_{PicoI}	P0: 0.9; <<0.01; 12	P0: n.s.	P0: n.s.	P0: n.s.	P0: n.s.	P0: 0.83; 0.04; 6
	P1: 0.82; <<0.01; 48	P1: 0.64; <<0.01; 24	P1: 0.53; <0.01; 24	P1: 0.6; <0.01; 24	P1: 0.65; <<0.01; 24	P1: n.s.
	P2: 0.36; <<0.01; 110	P2: n.s.				
	P3: -0.28; 0.05; 51	P3: n.s.	P3: -0.34; 0.05; 33	P3: n.s.	P3: n.s.	P3: n.s.
BV_{PicoII}	P0: -0.76; <0.01; 12	P0: n.s.	P0: n.s.	P0: n.s.	P0: 1; <<0.01; 6	P0: 0.94; <0.01; 6
	P1: 0.6; <<0.01; 48	P1: 0.54; <0.01; 24	P1: 0.4; 0.05; 24	P1: 0.58; <0.01; 24	P1: 0.63; <0.01; 24	P1: n.s.
	P2: n.s.	P2: n.s.	P2: n.s.	P2: 0.54; <<0.01; 42	P2: n.s.	P2: n.s.
	P3: 0.36; 0.01; 51	P3: 0.46; <0.01; 38	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.
BV_{PicoIII}	P0: n.s.	P0: n.s.	P0: n.s.	P0: n.s.	P0: n.s.	P0: n.s.
	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.
	P2: 0.6; <<0.01; 112	P2: n.s.	P2: 0.3; 0.05; 42	P2: 0.42; <0.01; 42	P2: 0.7; <<0.01; 42	P2: n.s.
	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.
BV_{NanoI}	P0: n.s.	P0: n.s.	P0: n.s.	P0: n.s.	P0: 1; <<0.01; 6	P0: 0.94; <0.01; 6
	P1: 0.45; <<0.01; 48	P1: n.s.	P1: 0.4; 0.05; 24	P1: n.s.	P1: 0.4; 0.05; 24	P1: n.s.
	P2: -0.53; <<0.01; 112	P2: n.s.	P2: 0.44; <0.01; 42	P2: 0.43; <0.01; 42	P2: n.s.	P2: -0.44; <0.01; 42
	P3: -0.35; 0.03; 51	P3: n.s.	P3: 0.41; 0.02; 33	P3: n.s.	P3: n.s.	P3: n.s.
BV_{NanoII}	P0: n.s.	P0: n.s.	P0: n.s.	P0: n.s.	P0: 0.81; 0.05; 6	P0: n.s.
	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.	P1: n.s.
	P2: -0.76; <<0.01; 112	P2: -0.37; 0.02; 42	P2: n.s.	P2: n.s.	P2: -0.46; <0.01; 42	P2: -0.34; 0.03; 42
	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.	P3: n.s.



15 Figure 1. A) Concentration of Chlorophyll *a* [$\mu\text{g L}^{-1}$], B) biovolume of nanophytoplankton
16 (Nano I and Nano II) [$\times 10^5 \mu\text{m}^3 \text{ml}^{-1}$] and C) biovolume of picophytoplankton
17 (*Synechococcus* spp., Pico I-III) [$\times 10^5 \mu\text{m}^3 \text{ml}^{-1}$] during the course of the experiment. Colours
18 and symbols indicate average $f\text{CO}_2$ [μatm] between t1-t43.

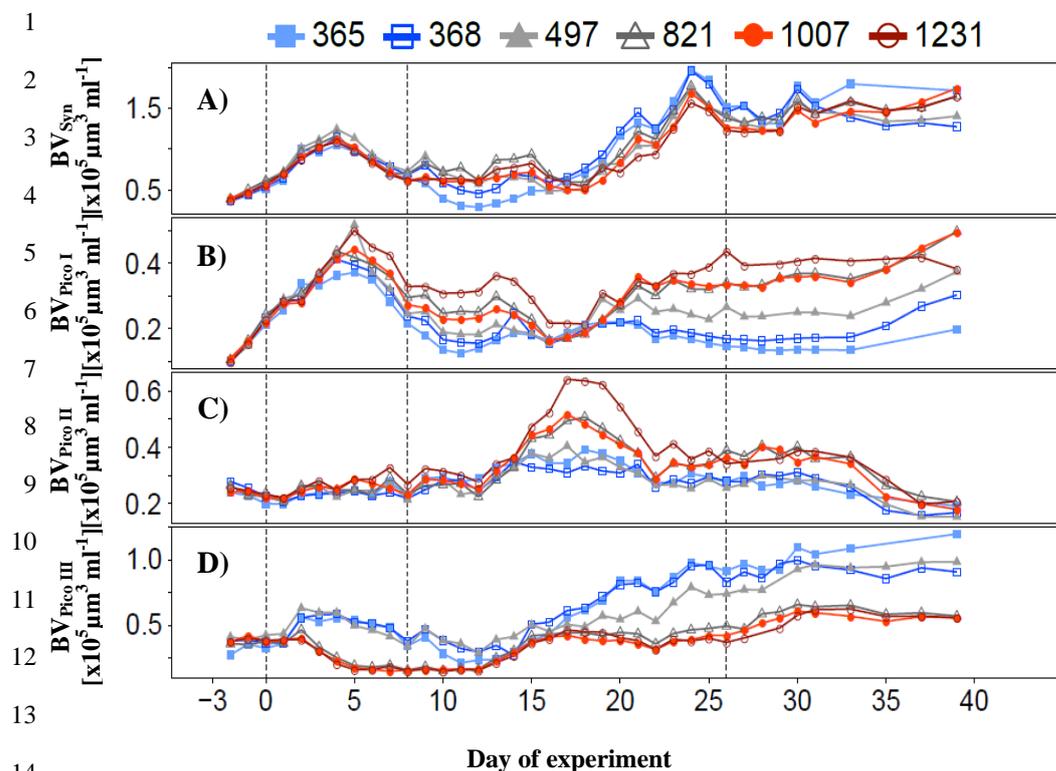
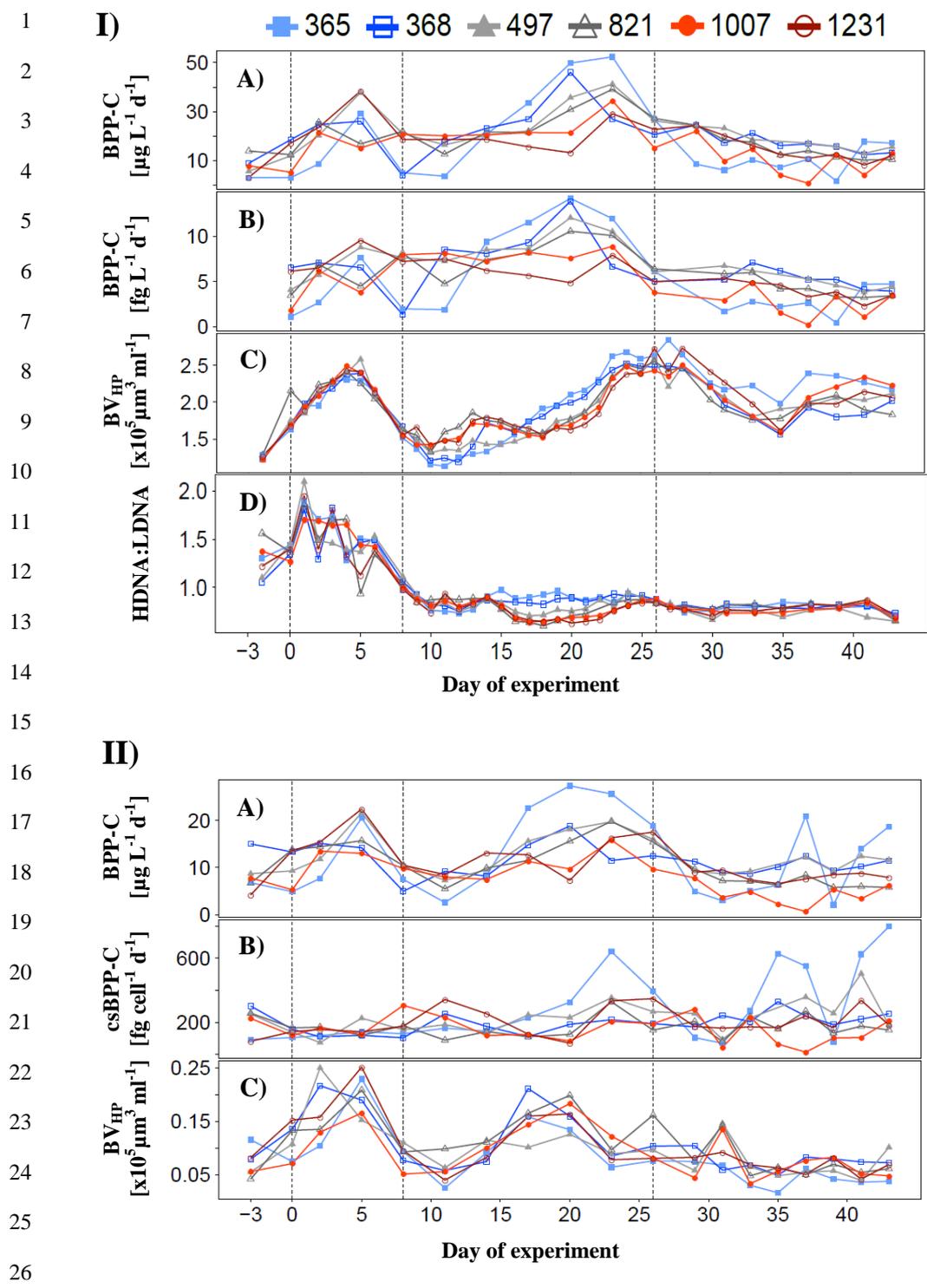


Figure 2. A) biovolume of *Synechococcus* spp. [$10^5 \mu\text{m}^3 \text{ml}^{-1}$] and B-D) biovolume of
 picoeukaryote groups I-III (Pico I-III) [$10^5 \mu\text{m}^3 \text{ml}^{-1}$] during the course of the experiment.
 Colours and symbols indicate average $f\text{CO}_2$ [μatm] between t1-t43.





1 Figure 3. A) Bacterial Protein Production (BPP-C) [$\mu\text{g L}^{-1} \text{d}^{-1}$], B) cell-specific Bacterial
2 Protein Production (csBPP-C) [$\text{fg cell}^{-1} \text{d}^{-1}$] and C) biovolume of heterotrophic prokaryotes
3 (BV_{HP}) [$\times 10^5 \mu\text{m}^3 \text{ml}^{-1}$] of size fractions I) 0.2-5.0 μm (free-living bacteria) and II) $>5.0 \mu\text{m}$
4 (particle-associated bacteria) during the course of the experiment. D) Ratio of high versus low
5 nucleic acid stained prokaryotic heterotrophs (HDNA:LDNA), which made up free-living
6 BV_{HP} , revealed from flow cytometry. Colours and symbols indicate average $f\text{CO}_2$ [μatm]
7 between t1-t43.

8

9

10

11

12

13

14

15

16

17

18

19

20

21

22

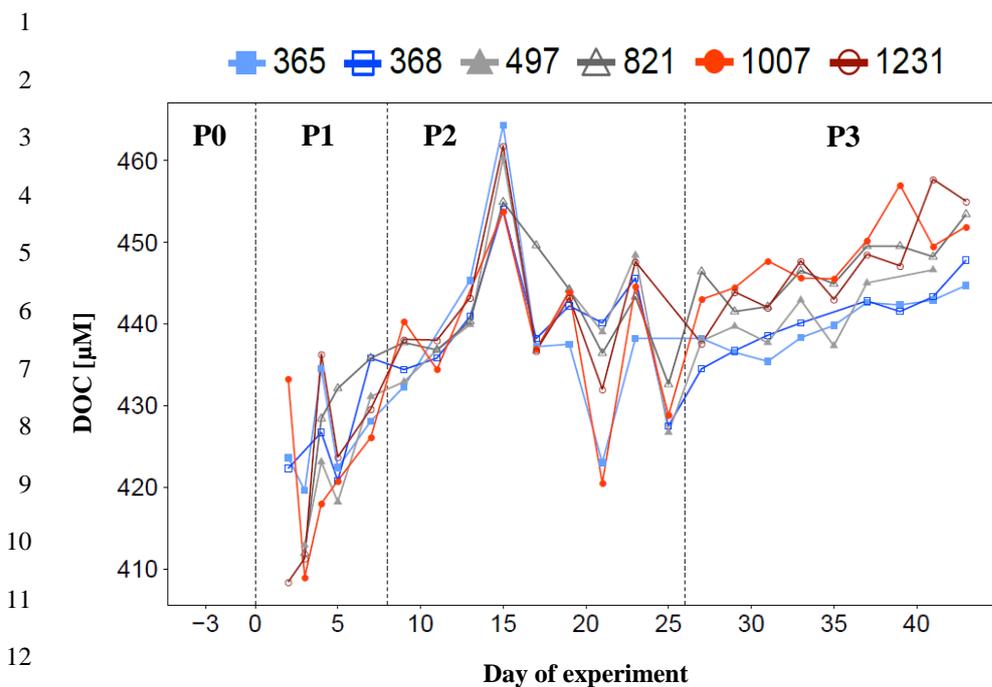
23

24

25

26

27



14 Figure 4. Concentration of dissolved organic carbon (DOC) [μM] during the course of the
15 experiment. During P3, DOC accumulated in the water column, thereby yielding significantly
16 higher concentrations at higher $f\text{CO}_2$ ($r_s=0.62$; $p<<0.01$; $n=51$). Colours and symbols indicate
17 average $f\text{CO}_2$ [μatm] between t1-t43.

18
19
20
21
22
23
24
25
26

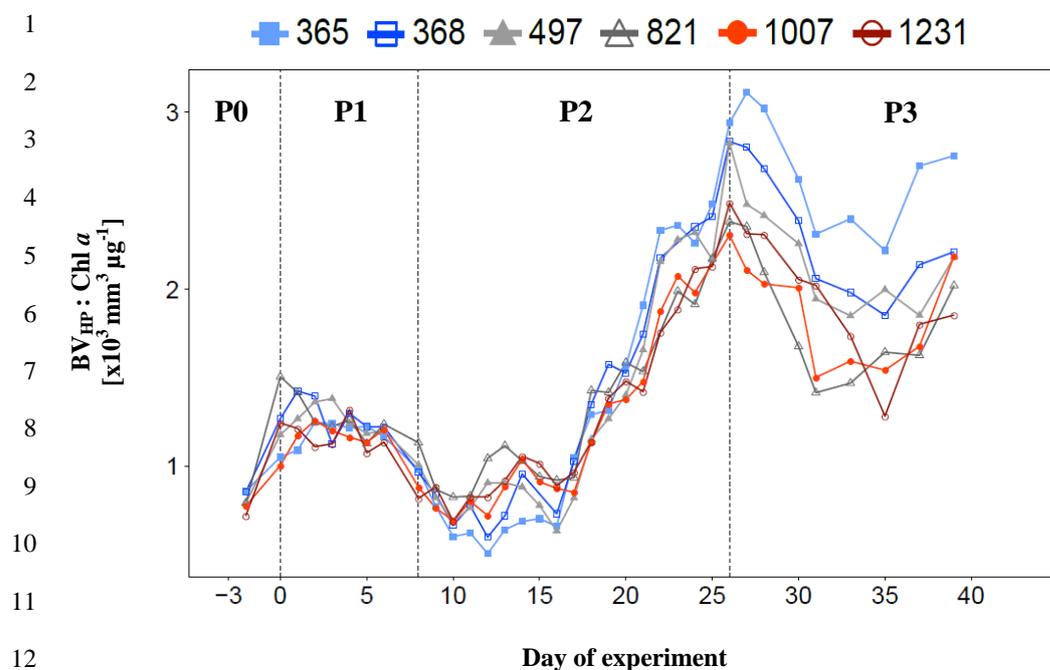


Figure 5. Standardization of heterotrophic prokaryotic biovolume to total Chl *a* ($BV_{HP} : Chl a$) during the course of the experiment. Colours and symbols indicate average fCO_2 [μatm]